

Model calculations of the migration of radionuclides from a repository for spent nuclear fuel

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SVENSK KÄRNBRÄNSLEFÖRSÖRJNING AB / AVDELNING KBS Swedish Nuclear Fuel Supply Co/Division KBS MAILING ADDRESS: SKBF/KBS, Box 5864, S-102 48 Stockholm, Sweden Telephone 08-67 95 40 MODEL CALCULATIONS OF THE MIGRATION OF RADIONUCLIDES FROM A REPOSITORY FOR SPENT NUCLEAR FUEL

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This report concerns a study which was conducted for SKBF/KBS. The conclusions and viewpoints presented in the report are those of the author(s) and do not necessarily coincide with those of the client.

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Stockholm maj 1983

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SUMMARY

Release and migration calculations for all important radionuclides which potentially could escape from a repository for spent nuclear fuel have been performed. The calculations have been done for a repository in fissured crystalline rock with models which account for the leaching of the nuclides from the fuel due to dissolution in the moving water and the further transport along the fissures of the rock. Radionuclide sorption retarding the nuclides is accounted for by the concept of matrix diffusion and sorption on the microfissure surfaces. Dispersion and channeling are modeled as Fickian dispersion. Chain decay is accounted for. Calculations have been performed for migration distances of 5, 50 and 100 m. Sorption data and solubility data for oxidizing as well as reducing conditions have been used.

Transport by sorption on colloidal matter in the water is also treated.

In the "central" case the leaching starts after 100.000 years due to the long life time of the canister, but calculations have also been made for a case of some defective canisters where the leaching starts after 100 years. The results are presented as release rates to the biosphere at the various times.

i

CONTENTS

1.	OBJECT	IVES	1
2.	CALCULA 2.1	ATION PRINCIPLES General description	2
	2.2 2.2.1 2.2.2 2.2.3	Activity release to the geosphere General Release in dissolved form Release in colloidal form	2 4 6
	2.3 2.3.1 2.3.2	Activity release to the biosphere Migration in dissolved form Migration in colloidal form	7 11
3.	INPUT P 3.1	ARAMETERS Common parameters	12
	3.2	Scenario specific parameters	17
4.	RESULTS 4.1	Summary of the calculations performed	18
	4.2 4.2.1 4.2.2 4.2.3 4.2.4	Activity release 1.0 $1/m^2$, year 50 m 0.3 $1/m^2$, year 100 m 0.1 $1/m^2$, year 100 m 0.03 $1/m^2$, year 5 m	20 21 49 61 85
5.	THE EFF AVAILAB 5.1 5.2 5.2.1 5.2.2 5.2.3 5.2.3 5.2.4 5.3	ECT OF 10% OF THE IODINE AND CESIUM LE IN PROMPTLY LEACHABLE FORM Background Release from a whole repository General Release to the geosphere of dissolved material Release to the biosphere of dissolved material Release to the biosphere of colloidally bound material Summary	97 98 98 100 100
6.	DISCUSSI	ION	101
7.	SYMBOLS	USED	103
8.	REFERENC	les .	104
APPENDIX A The Bateman equations APPENDIX B Conversion of computer program TRUCHN from concentration to activity units.			
APPENDIX C Steady state solution for a constant source			

1. OBJECTIVES

This study is a part of the Swedish KBS-3 project.

The aim of the study is to assess the release to the geosphere and to the biosphere of radioactive material from a repository for spent unreprocessed nuclear fuel.

2. CALCULATION PRINCIPLES

2.1 General description

The release to the biosphere from a repository for unreprocessed nuclear waste has been calculated for four scenarios with different water flow rates and migration lengths.

The release calculations have been performed in two steps.

- i The release from the immediate neighbourhood (near field) of a canister, giving a source concentration for the geosphere migration model.
- ii The release after passage through a geological medium.

The two transport mechanisms which have been modeled are with the radionuclides in dissolved form or adsorbed on colloidal particles.

Parameter variations that have been performed for some or all of the different scenarios are: different transport mechanisms, time for start of the leach process, reducingor oxidizing migration conditions and different values of dispersion coefficients.

2.2 Activity release to the geosphere.

2.2.1 General description

At a certain time the copper-cladding of a canister is assumed to be penetrated as a result of corrosion. Water will then reach the fuel elements and the ionizing radiation will create an oxidizing environment. The oxidizing species (notably hydrogen peroxide and peroxide radicals) will then diffuse out of the canister, through the clay barrier and into the geologic medium together with dissolved fuel constituents in, when so possible, oxidized form.

The rock, at the sites proposed for final disposal of the nuclear waste, contains reducing substances, such as Fe²⁺ which will limit the spreading of the oxidizing zone and lower the oxidation state of some of the migrating species (Neretnieks I,Aslund B, 1983).

As the fuel matrix dissolves its constituents will be released at a rate proportional to the dissolution rate of uranium in oxidized form (congruent dissolution). Eventually the amount of material set free from the matrix may exceed what can be completely dissolved either in oxidized or reduced form. The interaction between different solid phases affecting the release of a member of a decay chain is schematically

depicted in figure 1.

FIGURE 1 INTERACTION BETWEEN DIFFERENT SOLID PHASES IN THE LEACH PROCESS OF A FUEL ELEMENT IN CONTACT WITH WATER.



Further assumptions:

- i Homogenous distribution of the fuel constituents except for 10% of the cesium and the iodine which are assumed to be promptly dissolved with a subsequent spread in time of the release caused by a finite rate of replacement of the water in the clay buffer. See further Chapter 5.
- ii No coprecipitation.

2.2.2 Release of dissolved material

Release from the fuel matrix

For homogenously distributed nuclides, the release from the fuel matrix can be written as:

M_{i,mat} ·C_{solub,U,ox} ·Q_{eq,ox}

.,0

As $M_{U,tot,mat} \approx M_{U238,mat}$ and the halflife of U238 in comparison with all the important nuclides, is very long we could approximate $M_{i,Mat}/M_{U,tot,mat}$ by:

<u>^Mi</u>	٠	Decay	factor		
^M 238U					(2)

(1)

where the decay factor is the relative change in content of the nuclide i caused by its own decay and by decay of its parent nuclides .

Release from the oxidizing zone

The maximum release from the oxidizing zone of the substance i can be written as:

C_{solub,j,ox}•Q_{eq,ox} (grams/year,canister) (3)

The actual release from the oxidizing zone will be given by (3) provided that the substance exists as a separate solid phase in oxidized form or the rate of release from the fuel matrix plus formation rate of the substance by decay of parent nuclides co-present as separate oxidized phase is larger than the solubility limit given by (3) In that case the release of the isotope i of the substance j will be:

```
\frac{M_{i,ox}}{\Sigma M_{i,ox}} \cdot C_{solub,j,ox} \cdot Q_{eq,ox} \qquad (grams/year,canister) \qquad (4)
```

Otherwise it will be equal to the release from the fuel matrix plus the formation rate from decay of parent nuclides in the separate solid oxidized phase.

Release from the redox front

C_{solub,j,red}•Q_{eq,red}

The maximum release from the redox front of the substance j can be written as:

(grams/year,canister) (5)

The actual release from the redox front can, in analogy with that from the oxidizing zone, be calculated as the minimum of:

 $\frac{M_{i,red}}{\sum_{ji,red}} \cdot C_{solub,j,red} \cdot Q_{eq,red} \qquad (grams/year,canister)$ (6)

and the actual release rate from the oxidising zone plus the formation rate from decay of parent nuclides in the separate solid reduced phase.

2.2.3 Release in colloidal form

An estimate of the outflow of radionuclides, adsorbed on colloidal particles, suspended in the water flowing past a canister, has been based on the following premises.

i The adsorption process is assumed to be irreversible.

ii The amount of substance adsorbed is assumed to be given by:

(7)(grams/year, canister) $K_{d} \cdot C_{coll} \cdot Q_{eq} \cdot C_{i}$

iii The actual release from the near field is given as the maximum of the hypothetical release from the oxidizing zone or from the redox front.

2.3 Activity release to the biosphere

2.3.1 Migration in dissolved form.

The geological medium in which the migration is thought to take place is modelled as cubical blocks of porous rock matrix into which constituents of the surrounding solution diffuses. The dimensions of these blocks are the same as the mean distance between fractures in the rock modelled.

In the calculation of the diffusion into the rock matrix the cubical blocks are approximated with spheres with the same surface to volume ratio as the cubical blocks.

The flowing water is thought to be confined to the interstices between the blocks as the permeability in the rock matrix is very low.

The transport is modelled as one dimensional. The only transverse transport is into the rock matrix.

The transport mechanisms in the fractures will be convective transport and hydrodynamic dispersion modelled as Fickian diffusion with a dispersion coefficient D given by a Peclet number Pe:

$$Pe = \frac{U \cdot L}{D_{L}}$$
(8)

The equation describing the transport of the i-th member in a radionuclide chain will be:

$$\frac{\partial c_{f}^{i}}{\partial t} + u_{f} \frac{\partial c_{f}^{i}}{\partial z} - D_{L} \frac{\partial^{2} c_{f}^{i}}{\partial z^{2}} = \alpha N^{i} - \lambda^{i} c_{f}^{i} + \lambda^{i-1} c_{f}^{i-1}$$
(9)

$$\kappa^{i} \frac{\partial c_{p}^{i}}{\partial t} = D_{p} \varepsilon_{p} \frac{\partial^{2} c_{p}^{i}}{\partial r^{2}} + \frac{\beta}{r} \frac{\partial c_{p}^{i}}{\partial r} - \kappa^{i} \lambda^{i} c_{p}^{i} + \kappa^{i-1} \lambda^{i-1} c_{p}^{i-1}$$
(10)

$$N^{i} = D_{p} \varepsilon_{p} \frac{\partial c_{p}^{i}}{\partial r} r = interface$$
(11)

.

Spherical blocks of rock matrix with a radius = r_0 means:

$$\alpha = -3(1-\varepsilon_{f})/\varepsilon_{f}r_{0}$$
(12)

$$\beta = 2 \tag{13}$$

The inflow boundary condition is a concentration boundary with a concentration given by:

$$C_{i,in} = Release_{i,min} Q_{eq,min}$$
 (Bq/1) (14)

Where min stands for the limiting part of the near field model (Dissolution rate of the fuel matrix, solubility in an oxidizing environment or solubility in a reducing environment).

The outflow boundary condition used is

$$c \rightarrow 0 \text{ for } z \rightarrow \infty$$
 (15)

The flux at the end of the migration distance is calculated as

$$\Phi_{i,out} = Q_{eq} \left[C_{i,out} - \frac{D_{L} \cdot \varepsilon_{f}}{U_{0}} \cdot \frac{\partial C_{i}}{\partial x} \right]$$
(16)

The calculations for single nuclides with a decaying band release as source term have in most of the cases been made with the analytical code NUCDIF (Rasmuson A, neretnieks I, 1982)

All the other calculations have been made with the TRUCHN code (Rasmuson et. al., 1982) . As this code in its present form is limited to maximum three nuclides of a decay chain and the source term description has to be either constant concentration for one nuclide or decaying band release for up to three members of a nuclide chain an approximation of the more detailed source term model outlined in 2.2.2 has to be utilized.

It can for instance be shown that the release to the biosphere of shortlived daughter nuclides with reasonably high retention will solely consist of the material formed during the migration of its parent nuclide and the primary release of the daughter nuclide will be insignificant for the migration calculation. One further approximation that has to

be used is for the case with vastly differing decay constants for the different members of a decay chain. This in many cases forces the TRUCHN model to utilize so short timesteps that the calculation for the whole decay chain can not be followed all the way to the maximum release without an excessive cost in terms of computer resources. The results from the test calculations that have been made usually show that for the case with successively shorter halflives when descending the decay chain means that release of a daughter nuclide assymptotically approaches the value predicted by the reconcentration concept (Rasmuson et.al., 1982). This means that the ratio of activity release of a nuclide to that of its parent nuclide is inversely proportional to the ratio of their K_d-values provided that the migration distance in terms of halflives for the daughter nuclide is sufficiently long for the "equilibrium" to be established and that the halflife of the daughter nuclide is much shorter than the halflife of its parent nuclide.

Also when the release of a daughter nuclide, in grams per year, grossly exceeds the release of the parent nuclide the calculation has been divided in two steps.

- i the migration of the parent nuclide
- ii the migration of the daughter, nuclide and the rest of the decay chain originating from that nuclide.

To assess the relative importance of the primary release of radioactive material from a leaching canister for the release to the biosphere, a calculation has been made for all the nuclides as a constant concentration source term. See appendix C.

The result of this calculation is that for the parameter combinations used the primary source terms for 231Pa, 226Ra, 230Th, 229Th have small or no effect on the release to the biosphere. The release to the biosphere for these nuclides has been calculated according to the reconcentration concept.

2.3.2 Migration in colloidal form

The migration is modelled as pure transport with the velocity of the water.

The colloidal particles are assumed to be transported to the biosphere with the velocity of the flowing water. The release to the biosphere is assumed to be the same as the release to the geosphere.

3. INPUT PARAMETERS

3.1 Common parameters used

The choice of parameter values is discussed in the KBS-3 report.

- The canisters are assumed to be penetrated at 10⁵ years after the closure of the repository. (10² years for the cases with initially defective canister).
- The respective inventories at 10^2 and 10^5 years are shown in tables 1 and 2.
- The solubilities in different media are shown in table 3.
- The K_{d} -values used are shown in table 4.
- The characteristic data for the rock are shown in table 5.
- The water flowrates and their corresponding equivalent flowrates are shown in table 6.

TABLE 1. Inventory at t=10⁵ years.

Nuclide	t _{1/2} (years)	Bq/canister	g/canister
 Zr93	1.53E6	9.84E10	1.06E3
Тс99	2.14E5	5.70E11	9.12E2
1129	1.57E7	1.86E9	2.85E2
Sn126	1.00E5	2.38E10	2.27El
Cs135	2.95E6	1.97E10	5.92E2
Ra226	1.60E3	6.22E10	1.70E0
Th229	7.34E3	2.54E10	3.12
Th230	8.00E4	6.22E10	8.64El
Th232	1.41E10	9.32E4	2.30E1
Pa231	3.28E4	9.84E8	5.63E-1
U233	1.59E5	2.80E10	7.55El
U234	2.45E5	9.84E10	4.26E2
U235	7.04E8	1.24E9	1.55E4
U236	2.34E7	1.92E10	7.71E3
U238	4.47E9	1.66E10	1.33E6
Np237	2.14E6	7.77E10	2.98E3
Pu239	2.41E4	8.81E11	3.83E2
Pu242	3.76E5	1.24E11	8.55E2

Nuclide	t _{1/2} (years)	Bq/canister	g/canister
H3	1.23+1	1.19+11	3.33-4
C14	5.73+3	2.86+10	1.73-1
Co60	5.27	5.18+8	1.24-5
Ni59	7.50+4	2.03+11	6.79+1
Ni63	1.00+2	1.59+13	7.57
Sr90	2.88+1	3.78+14	7.40+1
Zr93	1.53+6	1.13+11	1.21+3
Nb94	2.03+4	1.72+10	2.48
Tc99	2.14+5	7.77+11	1.24+3
I129	1.57+7	1.86+9	2.85+2
Cs135	2.95+6	2.02+10	6.08+2
Cs137	3.02+1	6.22+14	1.94+2
Sn126	1.00+5	4.71+10	4.48+1
Sm151	8.70+1	9.32+12	9.25
Eu154	8.5	2.23+11	2.21-2
Ho166m	1.20+3	2.02+8	3.04-3
Ra226	1.60+3	1.40+6	3.82-5
Th230	8.00+4	6.73+7	9.35-2
Th229	7.34+3	6.22+4	7.63-6
Th232	1.41+10	6.73+1	1.66-2
Pa231	3.28+4	2.54+6	1.45-3
U233	1.59+5	1.04+7	2.91-2
U234	2.45+5	9.32+10	4.04+2
U235	7.04+8	7.25+8	9.06+3
U236	2.34+7	1.40+10	5.84+3
U238	4.47+9	1.66+10	1.33+6
Np237	2.14+6	2.64+10	1.01+3
Pu238	8.77+1	8.81+13	1.39+2
Pu239	2.41+4	1.50+13	6.53+3
Pu240	6.57+3	1.97+13	2.35+3
Pu242	3.76+5	1.50+11	1.03+3
Am241	4.32	2.64+14	2.08+3
Am242m	1.52+2	3.94+11	1.09
Am243	7.37+3	1.71+12	2.31+2
Cm245	8.5+3	1.81+10	2.85
Cm246	4.70+3	4.56+9	3.98-1

TABLE 2. Inventory at $t=10^2$ years.

TABLE 3. Solubility data

Solubility (g/l)			
Oxidizing environment	Reducing environment		
0 36(0 95)*	10^{-5} (10 ⁻³)*8.10 ⁻² **		
e.10~6	3.10-3		
8,10	е.10 ⁻ б		
0.63	8-10 A-10 ⁻⁷		
4.10-7	4.10		
$6 \cdot 10^{-3}$	6·10 ⁻⁵		
High	2.10-/		
	Solubility (g/1) Oxidizing environment 0.36(0.95)* 8.10 ⁻⁶ 0.63 4.10 ⁻⁷ 6.10 ⁻³ High		

- * Values used for the cases with the water flowrate= 1.0 $1/m^2$, year and 0.3 $1/m^2$, year.
- ** Value used for some calculations on the 1.0 l/m²,year
 case.
- TABLE 4. Adsorption data

Substance	Mass absorption koefficient $K_d(m^3/kg)$
Co,Ni	0.2
Sr	0.04
Zr,Nb	4
Tc(ox/red)	0.0002/0.05
I	0
Cs	0.05
Ce,Nd,Eu,Ho,Sm	5
Ra	0.1
Th , Pa	5
U,Np(ox/red)	0.01/5
Pu	5 (0.5 if complex bound)
Am	5 " " " "

TABLE 5. Chara	acteristic data for th	e geologic medium
 Distance betwee Water flow por Porosity of the Effective diffective diffective diffective Feclet number 	een fractures = $2b = 5$ cosity $\varepsilon_{f} = 10^{-4}$ ne rock matrix $\varepsilon_{p} =$ tusivity in the rock m $U \cdot 1/D_{L} = Pe = 2$	m 0.002 atrix $D_p \epsilon_p =$ and 50
kg/m ³	of Colloidal particle	s C _{coll} =5·10 *
- Density of the	rock matrix $\rho_p = 2700$	kg/m ³
TABLE 6. Water flowr	flowrates and corresp ates	ponding equivalent
U _o 1/m ² , year	Q _{eq} ,ox	Q _{eq} ,red
	(l/canister,year)	(l/canister,year)
0.00	0.00	
0.03	0.32	4.5
1.U	0.00	15
1.0		45
T•0	1.5/(2.7)*	150

* The figures in parenthesis stands for the elevated flow around an initially defective canister which begins to leach at 100 years after the fuel has been taken out from the reactor. The temperature in the immediate vicinity of the canister is assumed to be 50 degrees centigrade above the ambient temperature. 3.2 Description of the different scenarios

The four main scenarios can be summarized as in table 7.

TABLE 7. Migration scenarios used

Water flow rate Migration distance (m) $U_0(1/m^2, year)$

0.1	100
1	50
0.03	5
0.3	100

4. RESULTS

4.1 Summary of the calculations performed

The parameters describing an individual set of calculations are:

The Peclet number Pe = 2 Pe = 50

The solubility of uranium in the reducing zone 10^{-5} g/l 10^{-3} g/l $8 \cdot 10^{-2}$ g/l

The solubility of uranium in the oxidizing zone 0.36 g/1 0.95 g/1

Migration conditions reducing conditions oxidizing conditions oxidizing conditions for uranium, reducing for the other nuclides

<u>Canister life time</u> 1·10⁵ years TABLE 8 Calculations performed

	2	Page
U ₀ =	1 1/m ² ,year L=50 m Release of dissolved material to the geosphere Release of dissolved material to the biosphere Release of colloidally bound material Release of dissolved material to the biosphere - comparison between Pe=2 and Pe=50	21 26 31 36
U ₀ =	<pre>1 l/m²,year L=100 m initially defective canister Release of dissolved material to the geosphere Release of dissolved material to the biosphere Release of colloidally bound material</pre>	38 42 45
U ₀ =	0.3 l/m ² ,year L=100 m Release of dissolved material to the geosphere Release of dissolved material to the biosphere Release of colloidally bound material	49 53 57
^U 0 ⁼	0.1 1/m ² ,year L=100 m Release of dissolved material to the geosphere Release of dissolved material to the biosphere Release of colloidally bound material	61 65 69
^U 0 ⁼	0.1 1/m ² , year L=100 m initially defective canister Release of dissolved material to the geosphere Release of dissolved material to the biosphere Release of colloidally bound material	73 77 81
U ₀ =	0.03 l/m ² ,year L=5 m Release of dissolved material to the geosphere Release of dissolved material to the biosphere Release of colloidally bound material	85 89 93

4.2 Activity release to the geosphere and to the biosphere

On the following pages the release of activity to the geosphere and to the biosphere is shown.

The curves for release to the biosphere are drawn only for the time span for which the calculations were made.

For instance calculations aborted because running out of CPU-time have been included where they clearly approach a limiting maximum value.

elease of dissolved material to the geosphere. $a=1 \ 1/m^2$, year Coolub II or $0.95 \ g/l$	
igure 2 The fission products except 99Tc	
igure 3 The natural decay chains and 99 Tc	
in oxidized form	
igure 4 The natural decay chains and 99Tc in reduced form.	

- C_{solub},U'red^{=8E-2} g/l Figure 5 The natural decay chains and 99Tc in reduced form.
- C_{solub,U,red}=1E-3 g/1
- TABLE 9. Existence of solid phase (years after start of leaching)

Substance	Oxidized	Reduced(8E-2g/1,U)	Reduced(1E-3g/1,U)
Тс		0-1E6	0-1E6
Th	0-2E8	-	5E4-1E7
Νр	-	0-2E6	0-2E6
U	0-9E5		0-8E6
Pu	0-3E6		-









```
Release of dissolved material to the biosphere.
U<sub>0</sub>=1 1/m<sup>2</sup>,year C<sub>solub</sub>,U,ox= 0.95 g/1
Migration distance = 50 m.
Figure 6 The fission product
Figure 7 The natural decay chains. Oxidizing migration
conditions
Figure 8 The natural decay chains. Reducing migration
conditions. C<sub>solub</sub>,U,red=8E-2 g/1
Figure 9 The natural decay chains. Reducing migration
conditions. C<sub>solub</sub>,U,red=1E-3 g/1
```









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Release of colloidally bound material $U_0 = 1 \ 1/m^2$, year $C_{solub}, U, ox^{=} 0.95 \ g/1$

- Figure 10 The fission products except 99Tc
- Figure 11 The natural decay chains and 99Tc in oxidized form
- Figure 12 The natural decay chains and 99Tc in reduced form. C_{solub},U,red = 8E-2 g/1
- Figure 13 The natural decay chains and 99Tc in reduced form. C_{solub},U,red⁼ 1E-3 g/1








Comparison between Pe=2 and Pe=50. Release of dissolved material to the biosphere. $U_0=1 \ 1/m^2$, year $C_{solub,U,ox}=0.95$ g/l. Reducing migration conditions

Figure 14 Release of 93Zr, 99Tc and 135Cs.

Comments

The larger the retention in the geologic medium the larger the effect of a lowered dispersivity. The comparison between 99Tc and 135Cs, which have the same retention, shows that the effect of an early arrival becomes increasingly important for nuclides with shorter halflife.



Release of	dissolved	material	to	the	geosphere.
U _o =1 1/m ² , Initially o	year C _{sol} defective	ub,U,ox ^{=0.} canister	,9 5 (g/l.	

- Figure 15 The fission products except 99Tc
- Figure 16 The natural decay chains and 99Tc in oxidized form
- Figure 17 The natural decay chains and 99Tc in reduced form. $C_{solub,U,red}=1\cdot10^{-3}$ g/1.
- TABLE 10. Existence times for solid phase (years after start of leaching).

Substance	The oxidizing zone	The redox front
Тс		0-2E6
Th	0-6E7	2E4-2E7
Np	-	0-2E6
U	0-6E5	0-1E7
Pu	0-2E6	-







Release of dissolved material to the biosphere. U₀=1 1/m²,year C_{solub},U,ox=0.95 g/1. Initially defective canister, migration distance=50 m. Figure 18 The fission products Figure 19 The natural decay chains. Reducing migration conditions.

Figure 18 The curves for 14C and 126 Sn are the same as in figure 15 - release to the geosphere.





11/1

Release of colloidally bound material. $U_0=1 \ 1/m^2$, year $C_{solub,U,ox}=0.95 \ g/l$. Initially defective canister.

- Figure 20 The fission products except 99Tc
- Figure 21 The natural decay chains and 99Tc from the oxidizing zone
- Figure 22 The natural decay chains and 99Tc from the redox front







Release of U ₀ = 0.3 l/m	dissolved material to the geosphere. ² ,year C _{solub,U,ox} = 0.95 g/1.
Figure 23	The fission products except 99Tc
Figure 24	The natural decay chains and 99Tc from the
	oxidizing zone
Figure 25	The natural decay chains and 99Tc from the
	reducing zone, C _{solub} ,U,red ^{=8·10⁻²} g/1
TABLE 11.	Existence of solid phase (years after start leaching)

Substance	The oxidizing zone	The redox front
Тс		0-2E6
Th	0-6E8	-
Np	-	0-5E6
U	0-2E6	
Pu	0-3E6	-

of







Release U _o =0.3 length	of 1/m ² = 10	dissolved material to the biosphere. ,year C _{solub,U,ox} =0.95 g/l migration 0 m.
Figure	26	The fission products
Figure	27	The natural decay chains. Reducing migration
		conditions. C _{solub} , U, red=8.10 ⁻² g/1
Figure	28	Oxidizing migration conditions for uranium.
		Reducing for the other nuclides.
Comment	:	
Figure	52	The curve for 23311 is the sum of the curves

Figure 52 The curve for 233U is the sum of the curves from primary release of 233U and 233U originating from migrating 237Np.







Release of colloidally bound material. $U_0=0.3 \ 1/m^2$, year $C_{solub}, U, ox^{=0.95} \ g/l$

- Figure 29 The fission products except 99Tc
- Figure 30 The natural decay chains and 99 Tc from the oxidizing zone
- Figure 31 The natural decay chains and 99Tc from the reducing zone.







Release of dissolved material to the geosphere. $U_0 = 0.1 \ 1/m^2$, year $C_{solub, U, ox} = 0.36 \ g/1$.

Figure 32 The fission products except 99Tc Figure 33 The natural decay chains and 99Tc in oxidized form Figure 34 The natural decay chains and 99Tc in reduced form $C_{solub,U,red}=1\cdot10^{-5}$ g/1

Comments:

Pu and Th will initially be left behind when the fuel matrix dissolves. U, Np and Tc will precipitate at the redox front and the production of Th from solid material at the redox front will soon exceed the solubility limit so a solid Th-phase will evolve there too.

TABLE 12.	Existence	of	solid	phase	(years	after	start	of
	release)							

Substance	In oxidized form	At the redox front
Тс	-	0-2E6
Th	0-3E9	3E4-1E9
Np	_	0-8E6
U	0-8E6	0-6E9
Pu	0-3E6	-







Release of dissolved material to the biosphere. $U_0=0.1 \ 1/m^2$, year $C_{solub}, U, ox=0.36 \ g/1$ migration distance =100 m



<u> 6</u>6



Release rate to the biosphere (Bq/year)


Release of colloidally bound material. U₀=0.1 1/m²,year C_{solub},U,ox=0.36 g/1 C_{solub},U,red=1·10⁻⁵ g/1. Figure 38 The fission products except 99Tc Figure 39 The natural decay chains and 99Tc in oxidized

form

Figure 40 The natural decay chains and 99Tc in reduced form.







Release of dissolved material to the geosphere. $U_0 = 0.1 \ l/m^2$, year $C_{solub}, U, ox^{=0.36} \ g/1$ initially defective canister.

0-3E6

Pu

Figure 41	The fission products exce	ept 99Tc
Figure 42	The natural decay chains a	and 99Tc from the
	oxidizing zone	
Figure 43	The natural decay chains a	and 99Tc from the redox
	front. C _{solub} ,U,red ^{=1·10⁻²}	5 g/l
TABLE 13.	Existence of solid phase leaching)	(years after start of
Substance	The oxidizing zone	The redox front
Тс	_	0-2E6
Th	0-2E9	2E4-2E9
Np	-	0-8E6
U	0-5E6	0-6E9







Release of dissolved material to the biosphere. U_o=0.1 1/m²,year C_{solub},U,ox=0.36 g/1 Initially defective canister, migration distance=100 m. Figure 44 The fission products Figure 45 The natural decay chains. Oxidizing migration conditions Figure 46 The natural decay chains, Reducing migration conditions. C_{solub},U,red=1·10⁻⁵ g/1





Release rate to the biosphere (Bq/year)



Release of colloidally bound material. U_o=0.1 1/m²,year C_{solub,U,ox}= 0.36 g/l Initially defective canister. Figure 47 The fission products except 99Tc Figure 48 The natural decay chains and 99Tc from the oxidizing zone Figure 49 The natural decay chains and 99Tc from the redox front







Release of	dissolved material to the geosphere.		
U ₀ =0.03 1/1	m ² ,year C _{solub} ,U,ox ⁼ 0.36	g/l	
Figure 50	The fission products exce	рt 99Тс	
Figure 51	The natural decay chains oxidizing zone	and 99Tc from the	
Figure 52	The natural decay chains	and 99Tc from the redox	
	<pre>front, C_{solub},U,red^{=1.10⁻}</pre>	⁵ g/l.	
TABLE 14. Existence of solid phase (years after start of		(years after start of	
	leaching)		
Substance	The oxidizing zone	The redox front	
Тс	-	0-2E6	
Th	0-8E9	1E4-1E10	
Np	-	0-1E7	
U	0-1E7	0-1E10	
Pu	0-3E6	-	







Release of dissolved material to the biosphere. $U_0=0.03 \ 1/m^2$, year $C_{solub}, U, ox^{=0.36}$ g/l migration distance = 5 m.

Figure 53 The fission products

- Figure 54 The natural decay chains. Oxidizing migration conditions
- Figure 55 The natural decay chains. Reducing migration conditions. C_{solub}, U, red^{=1·10⁻⁵} g/1.





Release rate to the biosphere (Bq/year)



Release of colloidally bound material. $U_0=0.03 \ 1/m^2$, year $C_{solub,U,ox} = 0.36 \ g/1$.

- Figure 56 The fission products except 99Tc
- Figure 57 The natural decay chains and 99Tc from the oxidizing zone
- Figure 58 The natural decay chains and 99Tc from the redox front. C_{solub,U,red}=1·10⁻⁵ g/1.







5. THE EFFECT OF 10% OF THE IODINE AND CESIUM AVAILABLE IN PROMPTLY LEACHABLE FORM 5.1 Background

Iodine and cesium have been shown to migrate to the grain boundaries in the fuel during the burnup period.

This means that these substances to some extent are more readily leachable.

In the calculations below it is assumed that 10 % of the total content of iodine and cesium promptly dissolves the moment water reaches the fuel.

The mechanism thought to spread out the release of the dissolved material in time is a finite rate of displacement of the water volume which initially contains the dissolved material. This volume is taken to be the 1.7 m³ of pore water residing in the clay buffer around a canister.

The primary release rate of promptly leachable material is shown in table 15.

TABLE 15 Primary Release rate of promptly leachable material (Bq/canister, year) (In parenthesis the duration of the release in years)

	0 1		_
1291	0.03	ann	3.5E4(5300)
	0.1	8.6E4(2100)	6.2E4(3000)
	1.0	3.0E5(630)	1.7E5(1100)
135Cs	0.03		3.7E5(5300)
	0.1	9.4E5(2100)	6.6E5(3000)
	1.0	3.2E6(630)	1.8E6(1100)
137Cs	0.03		
	0.1	3.0E10(2100)	
	1.0	9.9E10(630)	

Nuclide $U_0(1/m^2, year)$ at t=10² years at t=10⁵ years

5.2 Release from a whole repository

5.2.1 General

The collective effect of the release calculated above for a whole repository has been calculated on the following assumptions.

98

4400 canisters with one initially defective canister and an evenly distributed time for start of the leaching process for the rest of the canisters in the interval 10^5 to 10^6 years. This means that about 205 years will pass between each leach start.

For 0.03, 0.1 and 1.0 $1/m^2$, year groundwater flow rate this means that 26, 15 and 6 canisters will be leaching at the same time with the high rate calculated above.

5.2.2 Release to the geosphere of dissolved material

The release to the geosphere from the whole repository is shown in Table 16 below.

TABLE 16 Release to the geosphere of promptly leachable material (10%) from a repository consisting of 4400 canisters. (In parenthesis the maximum release from the other 90% residing in the fuel matrix calculated on the assumptions given elsewhere in this report)*.

Nuclide	⁰ 0	Release The initially	The other
	(l/m ² ,year)	defect. canister (Bq/year,canister)	canisters (Bq/year,4400 can.)
1291	0.03 0.1 1.0	8.6E4(3.9E2) 3.0E5(3.5E3)	9.1E5(7.0E5) 9.3E5(1.2E6) 1.0E6(9.1E6)
135Cs	0.03 0.1 1.0	9.4E5(4.3E3) 3.2E6(3.9E4)	9.6E6(6.6E6) 9.9E6(1.2E7) 1.1E7(9.1E7)
137Cs	0.03 0.1 1.0	3.0E10(1.3E8) 9.9E10(1.2E9)	

* Note: The different U₀ values also correspond to cases with different solubilities of uranium. The release rates shown above are not strictly additive. The effect on the maximum release rate to the biosphere will be very much dependent on the spread in time during the migration in the geosphere. 5.2.3 Release to the biosphere of dissolved material

100

For 129I the retention in the geological medium is very small and the figures given in table 16 are representative for the release to the biosphere.

For 135Cs the retention in the geological medium is sufficiently large to spread out the release from the promptly leachable material to such an extent as to make it indistinguishable from the release originating from the dissolution of the fuel matrix.

5.2.4 Release to the biosphere of colloidally bound material

The fraction of the release to the geosphere that will be transported to the biosphere by a colloidal mechanism will be:

 $2.5 \cdot 10^{-5}$ for Cs and 0 for I.

5.3 Summary

The effect of more readily leachable material on the release rate to the biosphere will be none for Cs and a maximum twofold increase in the maximum release rate to the biosphere of I.

The increase in release rate for I can be expected to be further diminished by the more probable successive degradation of the copper canisters.

6. DISCUSSION

The leach rate of the spent fuel is determined by the equivalent flow rate of water dissolving and transporting the main matrix constituent UO_2 . All other nuclides are released as the matrix dissolves away. The solubility of uranium in the water adjacent to the fuel is thus of prime importance. Due to the possible effects of alfa-radiolysis the water is assumed to be oxidizing and then uranium has a much higher solubility than in the naturally reducing waters in ferrous iron containing crystalline rock.

Some nuclides e.g. Pu and Th which have very low solubility will precipitate as the uranium dissolves away. Others e.g. Tc and Np which have very low solubility in reducing waters will precipitate at the redox front. For such nuclides the maximum concentration will be limited by their solubilities.

All dissolved nuclides will penetrate into the porous rock matrix by diffusion from the flowing water in the fissures. Most of the nuclides (except I and Sn) will sorb onto the inner surfaces of the pores or microfissures. This is by far the most important mechanism for retarding the nuclides in the present model. Compared to this sorption on the surfaces of the fissures with flowing water can be neglected.

Dispersion has been modelled as Fickian dispersion and high dispersivities have been used to ensure that early arrival of a portion of the nuclides is accounted for. (Sample calculations with small dispersion have shown to give smaller releases).

The treatment of colloidal transport by assuming that the nuclides are sorbed irreversibly on natural colloidal matter in the water is very conservative. If there were irreversible sorption mechanisms these would be active in the matrix also, effectively stopping a fraction of what now is modelled as reversibly sorbing nuclides.

In the present model calculations no account is taken of the retardation in the backfill. This is of little or no importance for the long lived canisters but would considerably retard some of the shorter lived nuclides released from a canister which fails initially.

Decrease of retardation by complex formation with fulvic and humic acids has been studied. 242 Pu was used as model nuclide. The release of this nuclide was increased but still did not become dominant. Other trivalent nuclides would be less important.

7.	SYMBOLS USED		
b	Half the fracture aperture	(m)	
С	Concentration	(g/l)	
D	Diffusivity in the pore water	(m ² /year)	
p D ₋	Dispersion coefficient in the flowing		
Ь	water	(m ² /year)	
К	Volume equilibrium constant =K _d p _n	(m ³ /m ³)	
K ,	Mass sorption coeffifient	(m ³ /kg)	
a L	Migration length	(m)	
М	Amount of a nuclide or substance	(kg)	
N	Molar flux at the fracture/solid		
	interface	(mol/m ² ,year)	
Ре	Peclet number = $u \cdot L/D_{\tau}$		
0	Equivalent water flow rate	(l/canister,year)	
~eq s	Spacing between fractures	(m)	
t	Time	(years)	
u	Linear velocity	(m/year)	
U_	Water flux in the undisturbed rock	(1/m ² ,year)	
U Z	Coordinate in the direction of		
	the water flow		
£	Porosity of the rock matrix		
p	Flow porosity		
	Density	(kg/m ³)	
Φ	Mass flux (g/	canister,year)	
λ	Decay constant	(year ⁻¹)	
Subscripts:	•		
coll	Colloidal particles		
f	The fissure, the flowing water		
i	Nuclide, isotope (Even used as a superscript)		
j	Substance		
ox	The oxidizing zone		
p	The rock matrix		
red	The reducing zone		
solub	Solubility		
8. REFERENCES

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104

The Bateman equations

The amounts Ni of the respective nuclides in an isolated decay chain 1 2 3 can be written as:

Nl = N0l * exp(-Ll *T)(1)

N2 = N01*L1/(L2-L1)*(exp(-L1*T)-exp(-L2*T))+N02*exp(-L2*T) (2)

$$N3 = N01*L1*L2*(exp(-L1*T)/(L2-L1)/(L3-L1)+ (3) + exp(-L2*T)/(L_1-L_2)/(L3-L2)+ + exp(-L3*T)/(L_1-L_3)/(L2-L3)+ + N02*L2/(L3-L2)*(expL-L2*T)-exp(-L3*T))+ + N03*exp(-L3*T)$$

Ni= amount of nuclide i (grams)
N0i= initial amount of nuclide i (grams)
T = time (years)
Li = decay factor for nuclide i (years⁻¹)

Conversion of computer program TRUCHN from concentration to activity units

As TRUCHN works with absolute accuracy, one should strive to get the different peaks in a radionuclide chain at about the same heights. For certain chains with long-lived parent nuclides, like U-238 (U-234) Th-230 Ra-226, it is then advantageous to work in activity units instead of the original formulation in concentration units in TRUCHN. The conversion is done in the following way.

The migration of the ith member of a radionuclide chain is, in concentration units, given by (Rasmuson and Neretnieks, 1982).

$$\frac{\partial c_{f}^{i}}{\partial t} + u_{f} \frac{\partial c_{f}^{i}}{\partial z} - D_{L} \frac{\partial^{2} c_{f}^{i}}{\partial z^{2}} = \alpha N^{i} - \lambda^{i} c_{f}^{i} + \lambda^{i-1} c_{f}^{i-1}$$
(1)

$$K^{i} \frac{\partial c^{i}_{p}}{\partial t} = D_{p} \varepsilon_{p} \frac{\partial^{2} c^{i}_{p}}{\partial r^{2}} + \frac{\beta}{r} \frac{\partial c^{i}_{p}}{\partial r} - K^{i} \lambda^{i} c^{i}_{p} + K^{i-1} \lambda^{i-1} c^{i-1}_{p}$$
(2)

$$N^{i} = D_{p} \varepsilon_{p} \frac{\partial c_{p}^{i}}{\partial r} r = interface$$
(3)

Spherical blocks of rock matrix with a radius = r_0 means:

$$\alpha = -3(1-\varepsilon_{f})/\varepsilon_{f} c_{0}$$
(4)

$$\beta = 2 \tag{5}$$

To obtain the analogue formulation in activity units we use:

$$a^{i} = \lambda^{i} c^{i}$$
 (6)

Substituting into equations (1)-(3) we obtain:

$$K^{i} \frac{\partial a_{p}^{i}}{\partial t} = D_{p} \varepsilon_{p} \left\{ \frac{\partial^{2} a_{p}^{i}}{\partial r^{2}} + \frac{\beta}{r} \frac{\partial a_{p}^{i}}{\partial r} \right\} - K^{i} \lambda^{i} a_{p}^{i} + K^{i-1} \lambda^{i} a_{p}^{i-1}$$
(7)

$$N^{i} = - D_{p} \varepsilon_{p} \left(\frac{\partial a_{p}^{i}}{\partial r} \right)_{r=interface}$$
(8)

$$\frac{\partial a_{f}^{i}}{\partial t} + u_{f}^{i} \frac{\partial a_{f}^{i}}{\partial z} - D_{L}^{i} \frac{\partial^{2} a_{f}^{i}}{\partial z^{2}} = -\alpha N^{i} - \lambda^{i} a_{f}^{i} + \lambda^{i} a_{f}^{i-1}$$
(9)

Equations (7)-(9) are the equivalents of (1)-(3) in activity units. The only difference is in the last terms of equations (7) and (8) (build-up due to the decay of a parent nuclide). They include the factor i instead of i-1.

The boundary conditions also have to be modified using equation (4). This is straight forward.

In the numerical code TRUCHN the modification (equations (7)-(9)) is done at the end of the main program HEART. The boundary conditions are appropriately modified in SUBROUTINE SURE.

Maximum release for an infinite source

Introduction:

To assess the importance of the primary source term to the geosphere migration calculation for a nuclide calculations have been performed on the stationary case with an infinite source i.e. a constant input concentration. The radioactive decay will establish a stationary concentration profile along the migration distance after a number of halflives for the nuclide studied.

An analytical expression has been derived for C/C_0 as a function of the migration parameters and the halflife of the nuclide (Neretnieks et.al. 1983). The mathematical expression can be written in the following way:

$$\frac{C}{C_0} = \exp\{L \cdot \left(\frac{Pe}{2L} - \left(\left(\frac{Pe}{2L}\right)^2 + \frac{\Psi}{D_L}\right)^{\frac{1}{2}}\right)\}$$
(1)

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$$\Psi = \lambda + \frac{\sqrt{\lambda} \cdot D_{e} \cdot r_{D}}{b}$$
(2)

Where:
$$\lambda = \text{the decay constant (years}^{-1})$$

 $D_e = \text{the effective diffusivity in the}$
rock matrix = $D_p \varepsilon_p (m^2/\text{year})$
 $r_d = \text{the retention factor } = K_d \text{ rock}$
 $2b = \text{fracture aperture (m)}$
 $Pe = \text{the Peclet number } = u_f L/D_L$
 $D_L = \text{the dispersion coefficient in the flow}$
 $direction (m^2/\text{year})$
 $L = \text{The migration length (m)}$
 $s = \text{fracture spacing (m)}$
 $\varepsilon_f = \text{flow porosity } (m^2/m^2)$

The case is derived for an infinite distance between fractures or in practical terms the fractures are so widely spaced that the concentration profiles from two fractures do not affect one another.

The penetration depths in the cases calculated for the KBS3 study ranges from millimeters to some centimeters for the nuclides notably affected by decay during the passage through the geologic medium. As this is a mere fraction of the 2.5 meters from the midline out to a fracture the diffusion process for the case with cubical blocks of rock matrix can effectively be modeled as diffusion into plane-parallel1 slabs.

To be able to compare with the solution of the numerical calculations, performed on qubical blocks of rock matrix with an edge length= S_{cube} and a fracture aperture=2b, the parameters b and D_L in (1) and(2) have to be modified.

For a medium with parallell fractures with a spacing=S and an aperture=2b, Snow (1968) has derived the expressions:

$$\sum_{a}^{3} b \propto \sqrt{s}$$

$$\epsilon_{f} \propto \frac{2b}{s}$$
(3)

First the parameter ψ which is common to all the cases studied, is listed in table 1.

and

Nuclide	t(1/2)(years)	𝕐 (m ^{−1})		
H3	1.23+1	5.64-2		
C14	5.73+3	1.21-4		
Co60	5.27	6.15+1		
Ni59	7.50+4	5.14-1		
Ni63	1.00+2	1.41+1		
Sr90	2.88+1	1.18+1		
Zr93	1.53+6	5.09-1		
Nh94	2.03+4	4.42		
Tc99	2.14+5	9.63-3/1.52-1		
I129	1.57+7	4.41-8		
Cs135	2.95+6	4.10-2		
Cs137	3.02+1	1.28+1		
Sn126	1.00+5	5.93-6		
Sm151	8.70+1	7.55+1		
Eu154	8.5	2.42+2		
Hol66m	1.2+3	6.65		
Ra226	1.6+3	2.49		
Th229	7.34+3	8.22		
Th230	8.00+4	2.49		
Th232	1.41+10	5.93-3		
Pa231	3.28+4	3.89		
U233	1.59+5	7.90-2/1.77		
U234	2.45+5	6.36-2/1.42		
U235	7.04+8	1.19-3/2.65-2		
U236	2.34+7	6.51-3/1.46-1		
U238	4.47+9	4.71-4/1.05-2		
Np237	2.14+6	4.81-1		
Pu238 red/ox(complex)	8.77+1	5.83+1/7.52+1(2.38+1)		
Pu239 "	2.41+4	3.51/4.54(1.43)		
Pu240 "	6.57+3	5.73/8.69(2.75)		
Pu242 "	3.76+5	8.85-1/1.14(3.61-1)		
Am241 (complex)	4.32+2	3.39+1(1.07+1)		
Am242m "	1.52+2	5.71+1(1.81+1)		
Am243 "	7.37+3	8.20(2.59)		
Cm245	8.5+3	7.64		
Cm246	4.7+3	1.03+1		
Common characteristics for the geologic medium:				
$D_e = 5 \cdot 10^{-14} m^2 / s$ $b_{cube} / 2 = 2.5 \cdot 10^{-4} m$ $s_{cube} = 5 m$				

The parameters special to each of the cases are shown in table 2. The C/C₀-values are shown in tables 3, 4 and 5. To get the same volume to surface ratio as for the case with cubical blocks we will have to use $S = S_{cube}/3$ (5) With a fixed Peclet number this means that $b=0.69 \cdot b_{cube}$ (6)

 $D_z = D_{z cube} / 2.08 \tag{7}$

TABLE 2. Parameters characteristic to each of the cases

Case:	U ₀ =1 1/1	m ,	U ₀ =0.3 1/m ² ,	U ₀ =0.1 1/m ² ,	U ₀ =0.03 1/m ² ,
	year Z=	50m	year Z=100 m	year Z=100 m	year Z=5 m
Pe	2	50	2	2	2
Pe/2 _L (m ⁻¹)	2·10 ⁻²	5•10 ⁻¹	1·10 ⁻²	1·10 ⁻²	2·10 ⁻¹
D _L (m ² /year)	1.2E2	4.81	7.26E1	2.40E1	3.61E-1
L(m)	50	50	100	100	5

TABLE 3. C/C _o for an infinite source. Values given for Pu and Am in parenthesis for $K_d=0.5 \text{ m}^3/\text{kg}$				
Nuclide	U _o =1 1/m ² ,year Z=50 m Pe=2	Pe=50		
3H	6.2E-1	5.6E-1		
14C	1.0	1.0		
60Co	7.6E-16	2.9E-68		
59Ni	8.9E-2	7.7E-3		
63Ni	9.5E-8	3.5E-96		
90Sr	4.1E-7	1.4E-25		
93Zr	9.0E-2	8.0E-3		
94Nb	1.8E-4	2.4E-13		
99Tc red/ox	9.1E-1/3.5E-1	9.0E-1/2.2E-1		
1291	1.0	1.0		
135Cs	7.0E-1	6.6E-1		
137Cs	2.1E-7	6.4E-27		
126Sn	1.0	1.0		
151Sm	1.6E-17	1.4E-76		
154Eu	3.9E-31	0		
166mHo	2.0E-5	1.3E-17		
226Ra	1.9E-3	6.8E-9		
229Th	5.4E-6	2.9E-20		
230Th	1.9E-3	6.8E-9		
232Th	9.4E-1	9.4E-1		
231Pa	3.2E-4	3.3E-12		
233U red/ox	5.3E-1/5.8E-3	4.5E-1/6.1E-7		
234U "	5.9E-1/1.1E-2	5.2E-1/6.7E-6		
235U "	9.9E-1/7.8E-1	9.9E-1/7.6E-1		
236U "	9.4E-1/3.6E-1	9.3E-1/2.3E-1		
238U "	1.0 /9.7E-1	1.0 /9.6E-1		
237Np "	9.8E-2	1.0E-2		
238Pu "	2.0E-15/1.7E-17(5.7E-10)	3.0E-66/2.1E-76(2.2E-39)		
239Pu "	5.0E-4/1.5E-4(1.1E-2)	2.3E-11/1.4E-13(6.2E-6)		
240Pu "	4.7E-5/3.8E-6(1.3E-3)	6.1E-16/5.2E-21(1.5E-9)		
242Pu "	3.3E-2/1.9E-2(1.5E-1)	3.6E-4/5.1E-5(3.0E-2)		
2 41 Am	7.7E-12(8.6E-7)	1.6E-48(5.0E-24)		
2 4 2 m Am	2.8E-15(9.8E-9)	1.8E-65(2.3E-33)		
2 4 3 Am	5.5E-6(1.6E-3)	3.1E-20(3.8E-9)		
245Cm	8.7E-6	2.6E-19		
246Cm	1.1E-6	1.9E-23		

TABLE	4.	c/c	for	an	infinite	source	val	ues	given	for	Ρu
	-	and	Am i	n p	parenthesi	s for	$K_{d} = 0.5$	_m ³ /	′kg	101	1 4

Nuclide	U ₀ =0.3 l/m ² , year Z=100 m Pe=2
3H	1.4E-1
14C	9.9E-1
60Co	2.9E-40
59Ni	5.7E-4
63Ni	2.0E-19
90Sr	8.3E-18
93Zr	5.9E-4
94Nb	5.1E-11
99Tc red/ox	5.9E-1/2.5E-2
1291	1.0
135Cs	2.1E-1
137Cs	1.6E-18
126Sn	1.0
151Sm	1.4E-44
154Eu	1.4E-79
166mHo	1.9E-13
226Ra	2.4E-8
229Th	6.5E-15
230Th	2.4E-8
232Th	7.1E-1
231Pa	2.4E-10
233U red/ox	8.7E-2/4.4E-7
234U "	1.2E-1/2.2E-6
235U "	9.2E-1/3.1E-1
236U "	6.9E-1/2.7E-2
238U "	9.7E-1/2.7E-1
237Np "	7.5E-4
238Pu "	3.3E-39/1.7E-44(3.7E-25)
239Pu "	7.5E-10/3.7E-11(2.1E-6)
240Pu "	1.7E-12/2.5-15(9.3E-9)
242Pu "	4.2E-5/9.4E-6(2.2E-3)
241 Am	5.7E-30(5.7E-17)
242 mAm	8.3E-39(5.6E-22)
243 Am	6.8E-15(1.7E-8)
245Cm	2.2E-14
246Cm	1.2E-16

TABLE 5.	C/C _o for an infinite source. parenthesis for Pu and Am are	Values given in for K _d =0.5 m ³ /kg
Nuclide	$U = 0.1 \ 1/m^2$.vear	$u_{a}=0.03 \ 1/m^{2}$, year
INGELICE	Z = 100 m	Z=5 m
	$P_{0} = 2$	Pe=2
L	re - 2	
ЗН	1.9E-2	3.0E-1
14C	9.8E-1	1.0
60Co	8.2E-70	1.2E-28
59Ni	1.2E-6	6.4E-3 7.2F-14
63N1	1.42-33	/•20 17
90Sr	9.5E-31	1.0E-12
93Zr 94Nb	1.2E-6 6.2F-19	6.7E-8
99Tc	2.9E-1/8.9E-4	7.5E-1/9.1E-2
1291	1.0	1.0
135Cs	3.9E-2	3.8E-1
137Cs	5.2E-32	3.1E-13
126Sn	1.0	1.0
151Sm	2.5E-77	1.1E-31
154Eu	0	0
166mHo	3.7E-23	1.3E-9
226Ra	2.7E-14	5.2E-6
229Th	1.0E-25	1.2E-10
230Th	2.7E - 14	5.2E-6
232Th	4.2E-1	8.3E-1
231Pa	8.8E-18	2.0E-7
233U	8.0E-3/4.3E-12	2.1E-1/4.0E-5
234U	1.4E-2/7.3E-11	2.7E - 1/1.3E - 4
2350	8.0E-1/8.5E-2 4.0E-1/1.0E-3	9.6E-1/0.0E-1 8.2E-1/9.7E-2
238U	9.1E-1/2.7E-1	9.8E-1/7.3E-1
237Np	4.9E-6	7.8E-3
0.2 0.0	F CD CO/2 CD 77/3 FE 42)	<pre>< ep. 20/1 20-31(6 30-19)</pre>
238PU 239Pu	5.6E-17/3.5E-19(6.7E-11)	4.5E-7/5.3E-8(1.2E-4)
240Pu	1.6E-2/2.0E-26(5.3E-15)	5.9E-9/5.9E-11(2.7E-6)
242Pu	1.2E-8/9.1E-10(1.2E-5)	1.0E-3/3.6E-4(1.7E-2)
241 Am	6.6E-52(2.7E-29)	2.4E-21(4.0E-12)
242mAm	2.8E-67(5.2E-38)	1.3E-27(1.1E-15)
243Am	1.1E-25(1.4E-14)	1.2E-1U(4.UE-6)
245Cm	8.5E-25	2.7E-10
246Cm	9.5E-29	6.7E-12

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