



**From soil clays to erosion processes**  
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# Clay erosion in safety assessment in Deep Geological Repositories (DGR)

Compacted bentonite will produce colloidal particles (Fissures; active fractures)

## Evaluation of clay loss

impact in barrier integrity and long term performance

Potential RN transport

## **Soil erosion:**

Unconfined state

Favorable conditions for colloidal release (rainwater physical impact; stream flow)

So, what can we learn from soil erosion studies to be applied in the SFA for DGR?



**The objective (as CIEMAT team suggested)  
Motivate new thinking/ideas to study clay erosion**

<http://paquita.blogspot.com.es/2014/09/pon-ton-oliva-carcavas-alpedrete-s-p.html>

# Summary

Soil aggregates stability: analogies to clay buffer erosion and colloid release control?

Physical-chemical and structural properties of different clay minerals:  
Clay systems erodibility

Factors influencing aggregate stability in soil clay systems: free oxides and organic matter

## Ideas



# 1. Soil aggregates stability: analogies to clay buffer erosion and colloid release control?

Type of treatment	Form of sample	Expression of the result	Authors
Wet sieving	3–5 mm	MWD*	Yoder (1936)
	<2 mm	% > 200 $\mu\text{m}$	Hénin <i>et al.</i> (195
	whole	<b>MWD:</b> mean weight diameter  Sum of mass fraction remaining on each sieve after sieving x mean aperture of the adjacent mesh : calculated range between 25 $\mu\text{m}$ -3.5 mm using the a set of six sieves	& I
	1–2		ser
	2–3.4		Ta
Raindrops or rainfall	1–2		y (
	4–5		
	2–9		
	5–8		
Ultrasonic dispersion	whole		
	4–5		ren
Immersion	4–5 mm	inter-aggregate pore volume	Grieve (1980)
	3–5 mm	qualitative	Emerson (1967)
Dry sieving	<4 mm	MWD	Kemper & Chepil (1965)

Tisdall and Oades (1982)

Macroaggregates  
(> 250 $\mu\text{m}$ )

Microaggregates  
(20-250  $\mu\text{m}$ )

Primary Particles  
(< 20 $\mu\text{m}$ )

Oades (1984)

Le Bissonais (1996)

Aggregate stability and assessment of soil crustability: I. Theory and methodology  
EJSS,47,425-437

# Aggregate Breakdown Mechanisms

Mechanism	Slaking	Breakdown by differential swelling	Breakdown by raindrop impact	Physico-chemical dispersion
Type of forces involved	Internal pressure by air entrapment during wetting	Internal pressure by clay differential swelling	External pressure by raindrop impact	Internal attractive forces between colloidal particles
Soil properties controlling the mechanism	Porosity, wettability, internal cohesion	Swelling potential, wetting conditions, cohesion	Wet cohesion (clay, organic matter, oxides) Elementary particles	Ionic status, clay mineralogy  Elementary particles Total
Resulting fragments	Microaggregates	Macro and microaggregates		
Intensity of the disaggregation	Large	Limited	Cumulative	

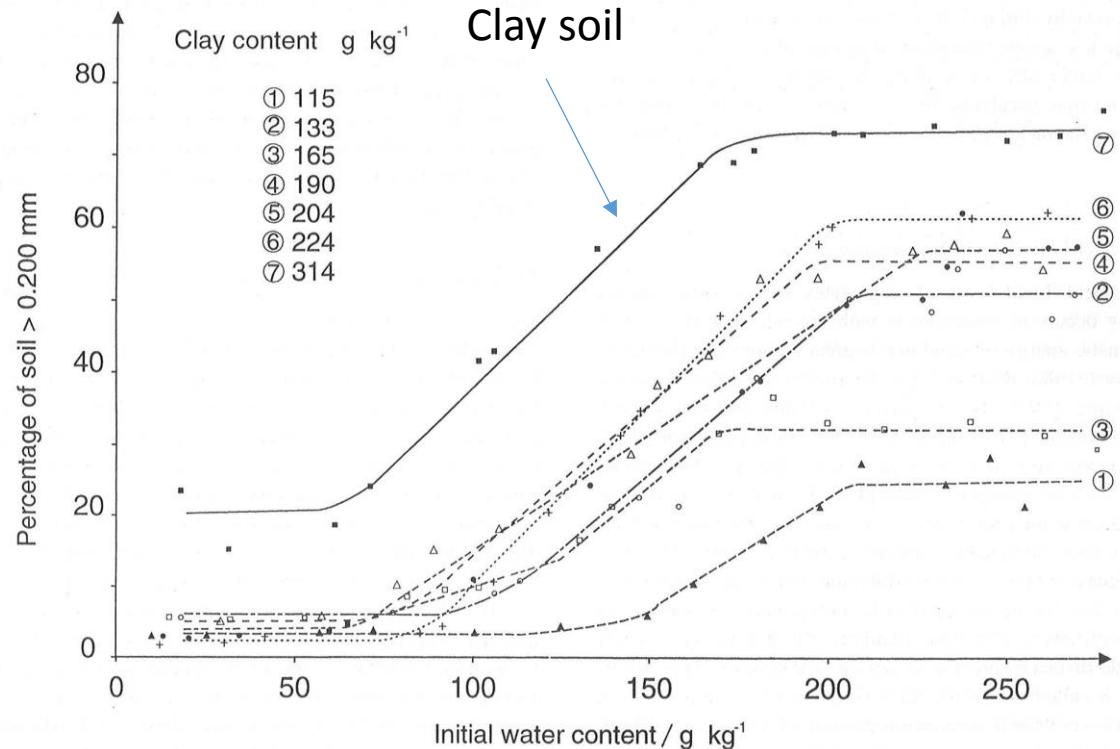
Slaking: rapid wetting

Differential swelling

Raindrop impact

Physico-chemical dispersion

Le Bissonais (1996)



# Analogies to clay buffer erosion and colloid release control?

## Colloid formation

### **In static systems: slaking?**

Rapid wetting of dry compacted bentonite: aggregates fragmented and dispersed by means of tensions generated by trapped air.

...the increase of clay particles in solution is initially rapid...

Differential swelling (micro-aggregates breakdown)

The bentonite loss in a static system is not to be continuous  
(crusting?: fissures clogging: self-healing)

**The dynamic system:** flow perturbations; physicochemical dispersion

Role of water chemistry; exchange complex, mixed materials ( free oxides, organic components)

### **Aggregate stability:**

porous filters < 1 $\mu$ m pore size do not allow the observation of particles release,

Indicators for micro (nano)-aggregate integrity measurements could be developed

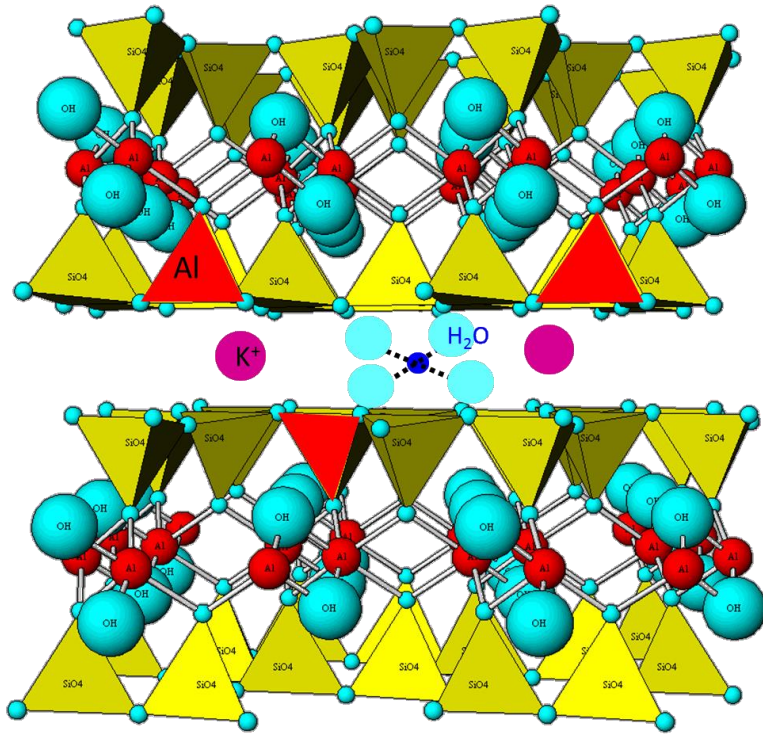
## 2. Physical-chemical and structural properties of different clay minerals: Clay systems erodibility



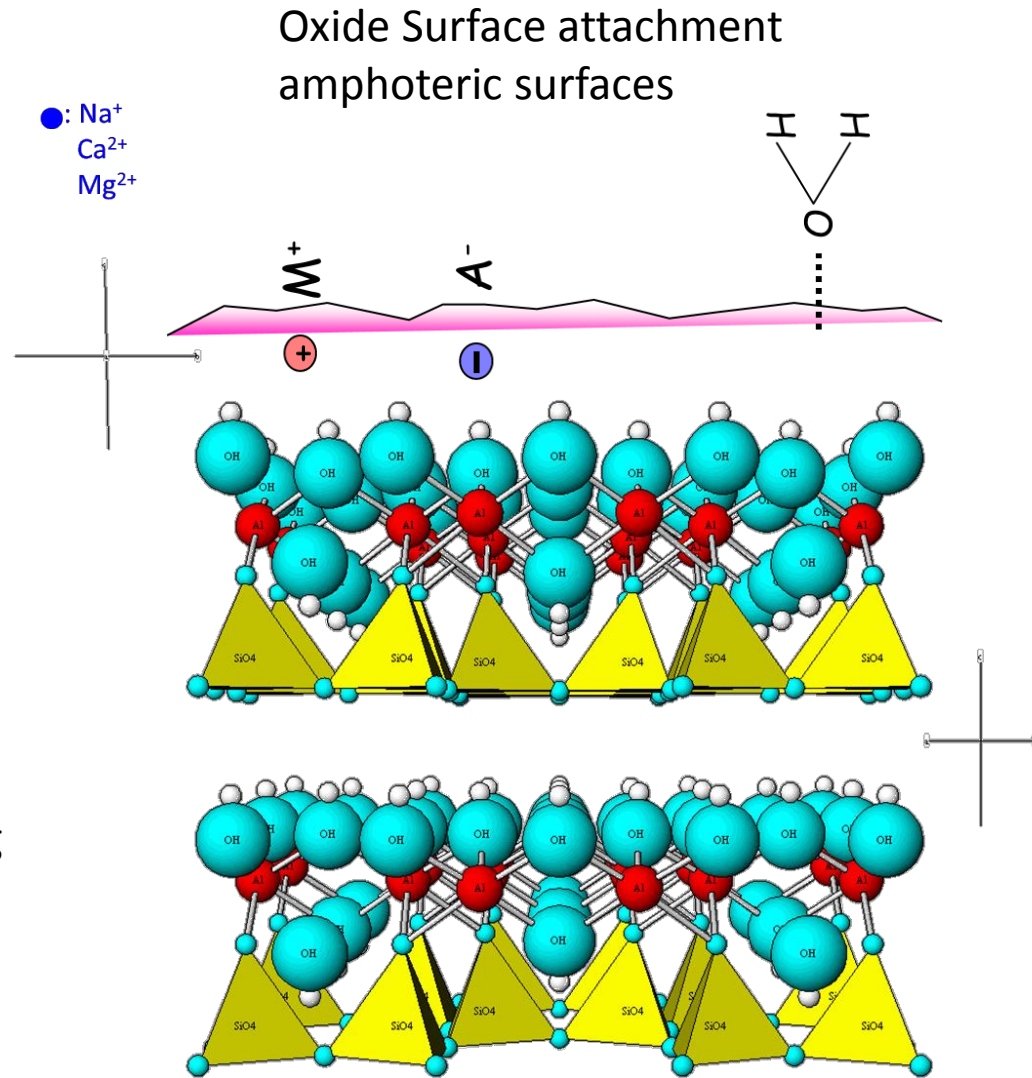


# CLAY MINERALS

## sheet 2:1 (illite $\text{K}^+$ , smectite $\bullet$ )



Soil organic matter (SOM)-M+-binding  
Organic cations binding  $\text{R-NH}_4^+$



# CLAY SYSTEM: ALSO A MATTER OF SIZE/SHAPE

## INTERACTION BETWEEN CLAY PARTICLES

NEVE YA'AR (MONTMORILLONITIC, 63% CLAY)

A Schematic structure  
of clay aggregate

B Scanning electron micrograph  
of clay tactoid



NETANY



TUNYAY



Mean weight diameter and total soil loss values for various studied soils (after Wakindiki and Ben-Hur, 2002)

Soil		Mean weight diameter		Total soil loss (kg m <sup>-2</sup> )
Location	Mineralogy	In fast-wetting test (mm)	In runoff (mm)	
		<b>SLAKING</b> →		<b>SEAL FORMATION UNDER RAIN FLOW</b> ←
Tunyai	Kaolinitic	2.80a <sup>a</sup>	0.12a	0.33a
Neve Ya'ar	Montmorillonitic	0.25b	0.03b	1.24b
Netanya	Montmorillonitic	0.31b	0.20c	1.14b
Molo	Non-phyllosilicate	0.84c	0.18c	0.75c
Njoro	Non-phyllosilicate	0.87c	0.21c	0.80c

<sup>a</sup> Different letters in a column indicate significant differences among soils,  $P \leq 0.05$ .





# CLAY SYSTEM: ERODIBILITY/(no oxide/SOM interaction)

## STABILITY OF COHESIVE MATERIALS AGAINST ERODITIVE FORCE OF FLOWING WATER

**Stability increased for 2:1 clay soil mixtures and for oriented kaolinite (1:1) matrix:**  
**function of clay content, previous water content. ERODIBILITY INCREASED WITH TEMPERATURE**

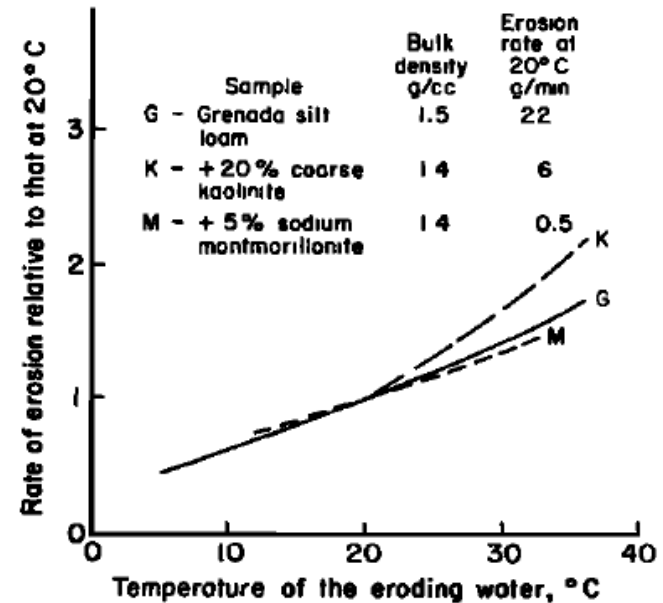
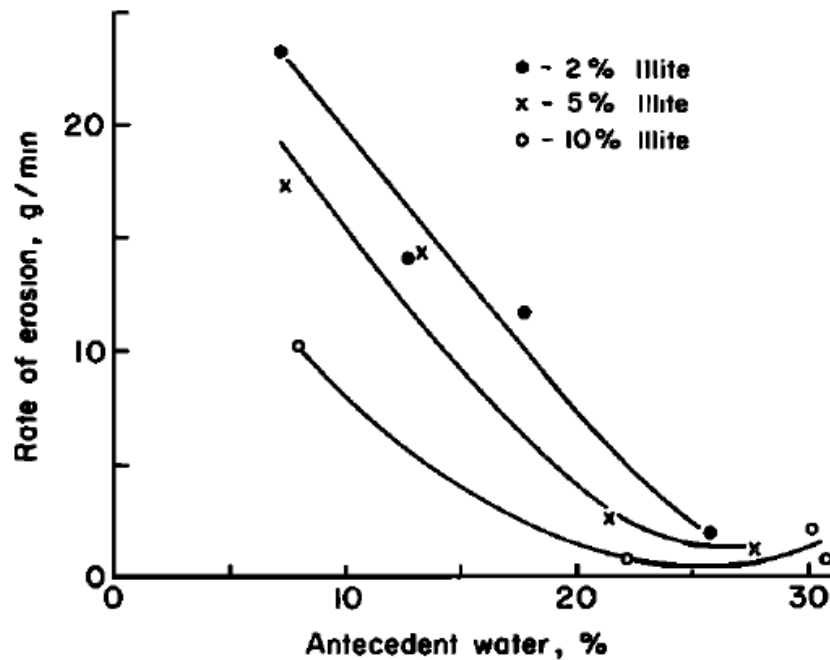


Fig. 1. Influence of temperature on erodibility of Grenada silt loam with admixtures, compacted and tested at 10% water.

E.H. GRISSINGER (1966)  
 Resistance Of Selected Clay Systems to  
 Erosion by Water  
 Water Resource Research, 2, 131-138

But: Erodibility (Interrill flow) is high for  
 smectitic soils and low for kaolinitic-oxidic  
 soils (Reichert et al., 2007)

### 3. Factors influencing aggregate stability in clay systems: free oxides and organic matter (SOM).

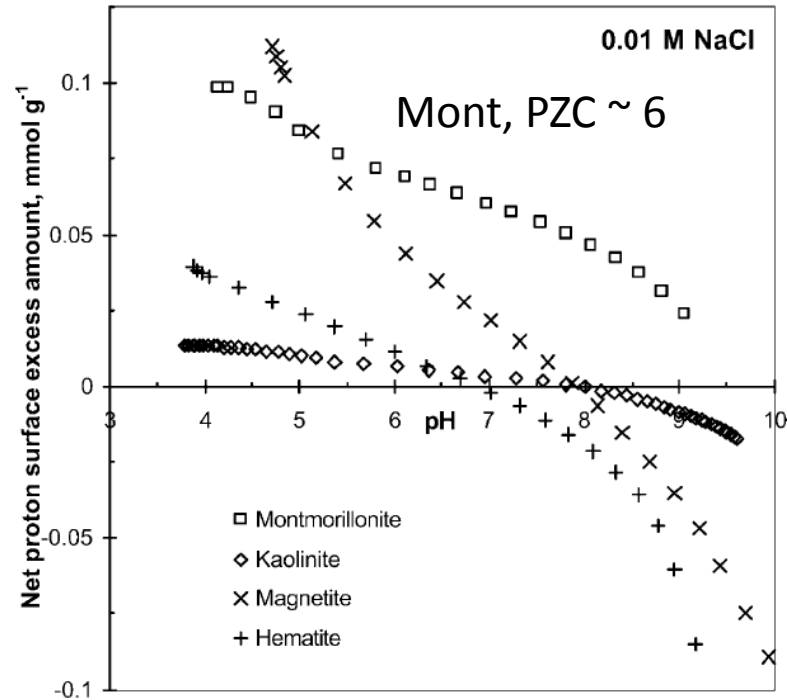


Fig. 1. pH-dependence of net proton surface excess amount determined by acid-base titration for clay minerals and metal oxides at 0.01 M ionic strength.

Proton Surface Excess acid- base titration  
PZC

Tombácz et al., 2004, Organic  
Geochemistry, 35, 257-267

**Small amounts of Al or Fe oxides improve flocculation of clay systems:** more effective for 1:1 systems at near the PZC of the oxides (7.2-8 Fe; 9.5 for Al).

(Goldberg and Glaubig, 1987; CCM, 35, 20-227)

In soils of mixed mineralogy (illite, smectite, kaolinite), **the removal of amorphous and crystalline oxides increases the clay dispersivity.**

But

**The removal of SOM decreased clay dispersivity.** Macro-aggregates are destroyed, SOM favors particle repulsion

Goldberg et al., 1990; Soil Science, 3, 588-593.



# CLAY SYSTEMS: organic matter (SOM) + Biomass interaction

## Macro-aggregate stabilization specific to 2:1 clay systems.

*J. Six et al. / Soil & Tillage Research 79 (2004) 7–31*

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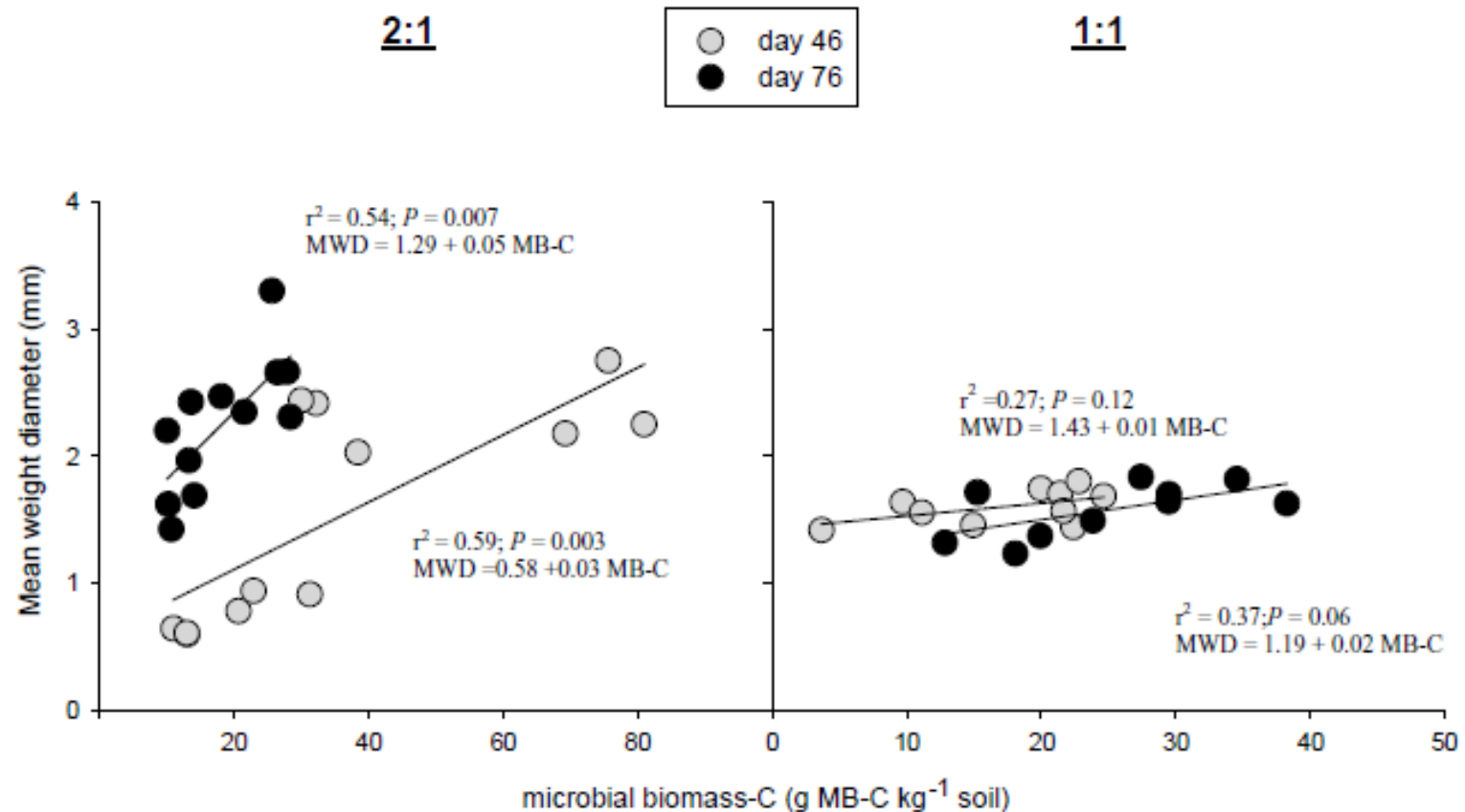


Fig. 7. Relationship between microbial biomass C (g MB-C kg<sup>-1</sup> soil) and aggregation (expressed in mean weight diameter) in a Mollisol dominated by 2:1 minerals and an Oxisol dominated by 1:1 minerals and oxides. Data adopted from Denef and Six (2003).

Illite, smectite selectively support/adsorb biopolymers related to microorganism activities

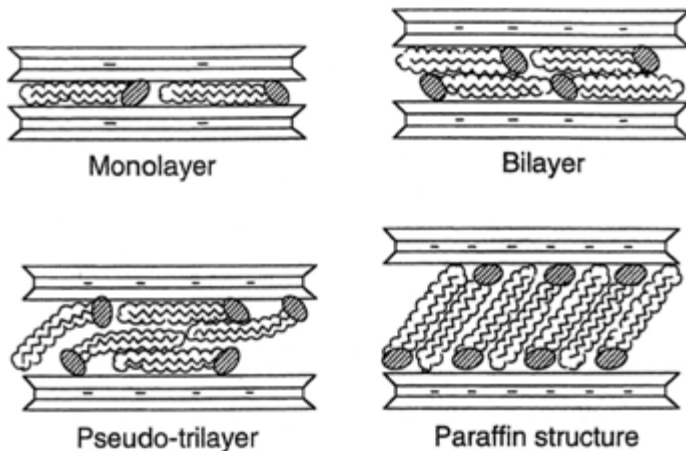
# Ideas

Materials design: Erodibility control.

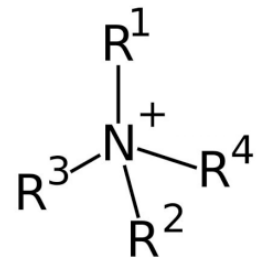
i.e.: **Iron oxides could control the PZC and favor aggregate stability at erosive interfaces.**  
Could be useful to apply oxide (i.e., magnetite or other iron oxide nanoparticles) to control erosion at the interfaces?

**Erosion measurements in field experiments: aggregate/particle detachment tracers:**  
organo-nano clays

i.e. Mentler et al., (2009). Organophylic clays as a tracer to determine erosion processes.  
Geophysical Research Abstracts, 11.: (detectable in 0.3  $\mu\text{g/L}$ )



Qutubuddin and Fu 2002









Thank you for your attention

Particle mediated transport linked to field erosion:

Cs at Fykushima

Antibiotics/pharma compounds

Potential applications of nanosized organoclay synthesis products

A choice for field or larger scale colloid transport evaluation?